



# Revisiting fundamental concepts in Mediterranean predictability: Liouville equation and Tailored bredes

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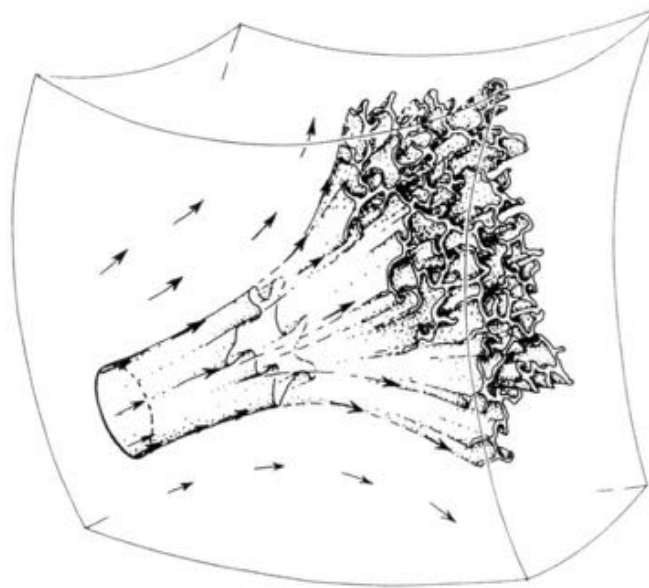
12th HyMeX Workshop, 20 - 23 May 2019, Split, Croatia

# Introduction

- Two sources of **uncertainty** in weather prediction: **initial state** and **physical processes**
- The mathematical framework to model uncertainty is **probability theory**
- **Uncertainty** is modelled through random variables **described by PDFs**. For multidimensional systems: multidimensional PDFs, representable over phase space

# Introduction

- Most characteristics of **atmospheric PDFs** are largely **unknown**. Certain topologies would challenge common interpretations of EPS.



Penrose, 1989

- The **perfect-model evolution** of a system described by a state vector of random variables fulfils the **Liouville equation**

# Objectives

- The purpose of this work is to explore some basic characteristics of simple solutions of the Liouville equation for low complexity systems
- Topological characteristics of the solutions will be analysed in detail
- Test the potential of LBVs to more efficiently initialize mesoscale EPS
- Investigate options to increase ensemble diversity and obtain a seamless scale representation compared to traditional BV

# PART I: Liouville equation

# General solution of the Liouville equation

- Liouville equation for the probability density function  $\rho$  given a dynamical system  $\dot{\mathbf{X}} = \Phi(\mathbf{X})$  :

$$\frac{\partial \rho(\mathbf{X}, t)}{\partial t} + \sum_{k=1}^N \Phi_k(\mathbf{X}, t) \frac{\partial \rho(\mathbf{X}, t)}{\partial X_k} = -\chi(\mathbf{X}, t) \rho(\mathbf{X}, t) \quad \chi = \sum_{k=1}^N \frac{\partial \Phi_k}{\partial X_k}$$

- If the **method of characteristics** can be used in the problem at stake, then the general solution is:

$$\rho(\mathbf{X}, t) = \rho_0(\boldsymbol{\alpha}) \exp\left(-\int_0^t \chi(\mathbf{X}(\boldsymbol{\alpha}, t'), t') dt'\right)$$

$\boldsymbol{\alpha}$  is the state vector at  $t = 0$

# Application to a Low dimensional barotropic model

- From barotropic vorticity equation under beta-plane approximation is:

$$\frac{\partial}{\partial t} \nabla^2 \psi = \frac{\partial \psi}{\partial y} \frac{\partial \nabla^2 \psi}{\partial x} - \frac{\partial \psi}{\partial x} \frac{\partial \nabla^2 \psi}{\partial y} - \beta \frac{\partial \psi}{\partial x}$$

- A solution consisting in a highly truncated Fourier series expansion for  $\nabla^2 \psi$  with three time dependent amplitudes ( $X_1$ ,  $X_2$  and  $X_3$ ) and phase angles linearly dependent on time ( $\theta_2$  and  $\theta_3$ ) is proposed (Peagle and Robl, 1977):

$$\nabla^2 \psi = X_1(t) \cos(l y) + X_2(t) \cos(k x + \theta_2) + 2X_3(t) \sin(k x + \theta_3) \sin(l y)$$

$$\theta_2 = \frac{\beta}{k} t$$

$$\theta_3 = \frac{\beta k}{k^2 + l^2} t$$

# Analytical solution of the system

$$X_1 = X_1^* \operatorname{dn}(h\tau + \phi, k_0^2)$$

$$X_2 = X_2^* \operatorname{sn}(h\tau + \phi, k_0^2)$$

$$X_3 = X_3^* \operatorname{cn}(h\tau + \phi, k_0^2)$$

$$\tau = \frac{\sin(\gamma t)}{\gamma}$$

dn, sn and cn are the Jacobi elliptic functions

If  $\beta$  (and  $\gamma$ ) are 0,  $\tau$  is t

$$k_0^2 = \frac{1 - \frac{C_2}{C_3} \frac{X_{30}^2}{X_{20}^2}}{1 - \frac{C_2}{C_1} \frac{X_{10}^2}{X_{20}^2}}$$

$$\phi = \operatorname{sn}^{-1} \left[ \frac{1}{\sqrt{1 - \frac{C_2}{C_3} \frac{X_{30}^2}{X_{20}^2}}}, k_0^2 \right]$$

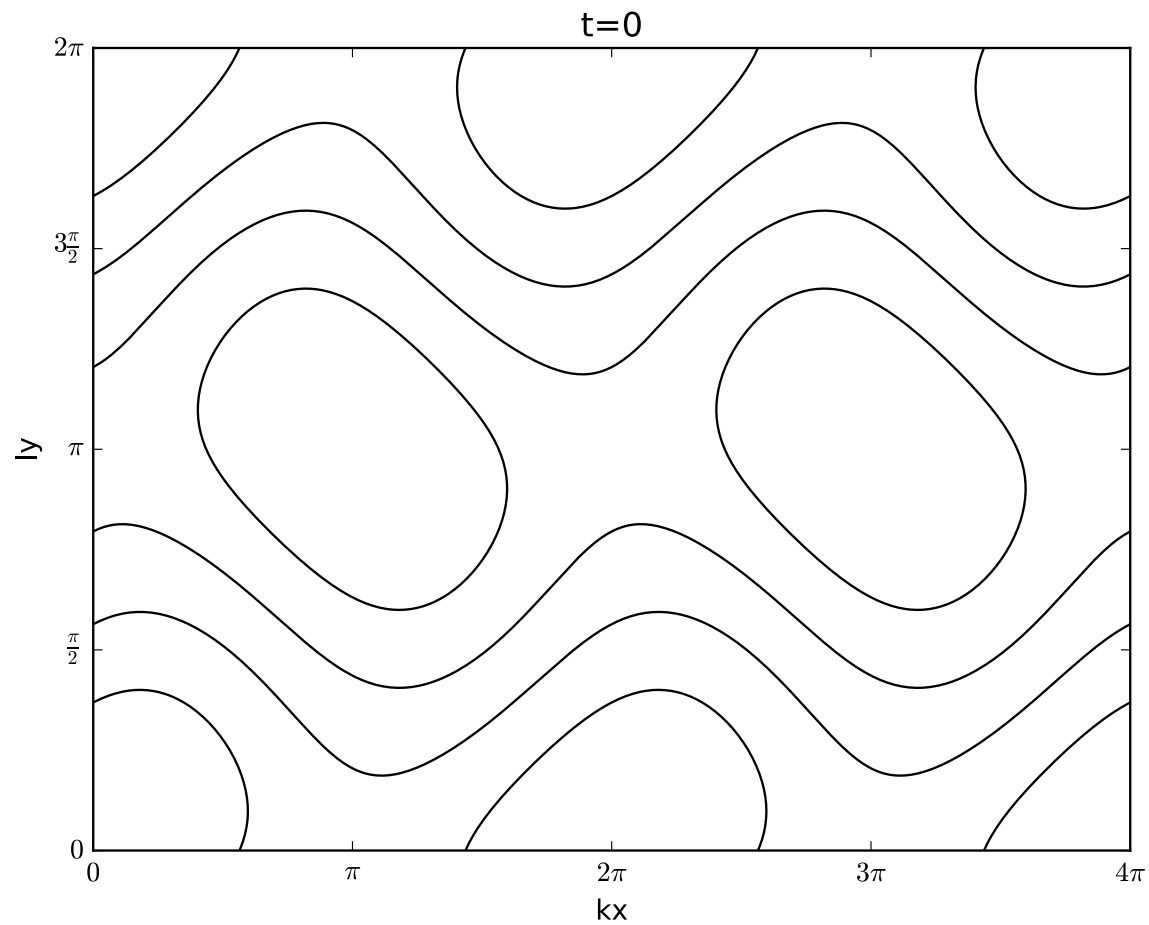
$$X_1^* = \frac{X_{10}}{\operatorname{dn}(\phi, k_0^2)}$$

$$X_2^* = \frac{X_{20}}{\operatorname{sn}(\phi, k_0^2)}$$

$$X_3^* = \frac{X_{30}}{\operatorname{cn}(\phi, k_0^2)}$$

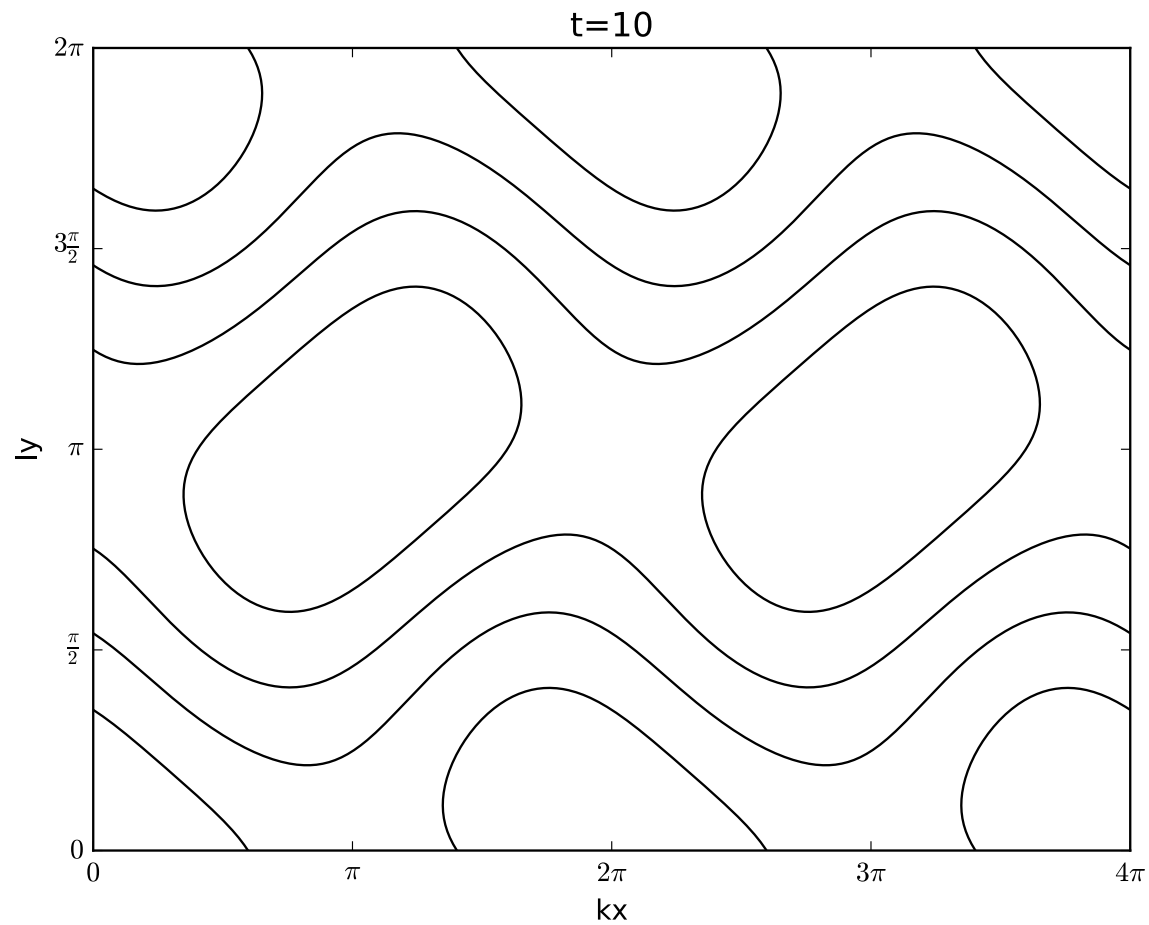
$$h = C_2 \frac{X_1^* X_3^*}{X_2^*}$$

# Analytical solution of the system



$$\begin{aligned} X_{10} &= 0.12 \\ X_{20} &= 0.24 \\ X_{30} &= 0.10 \\ \alpha &= 2 \\ \beta &= 0 \end{aligned}$$

# Analytical solution of the system



$$X_{10} = 0.12$$

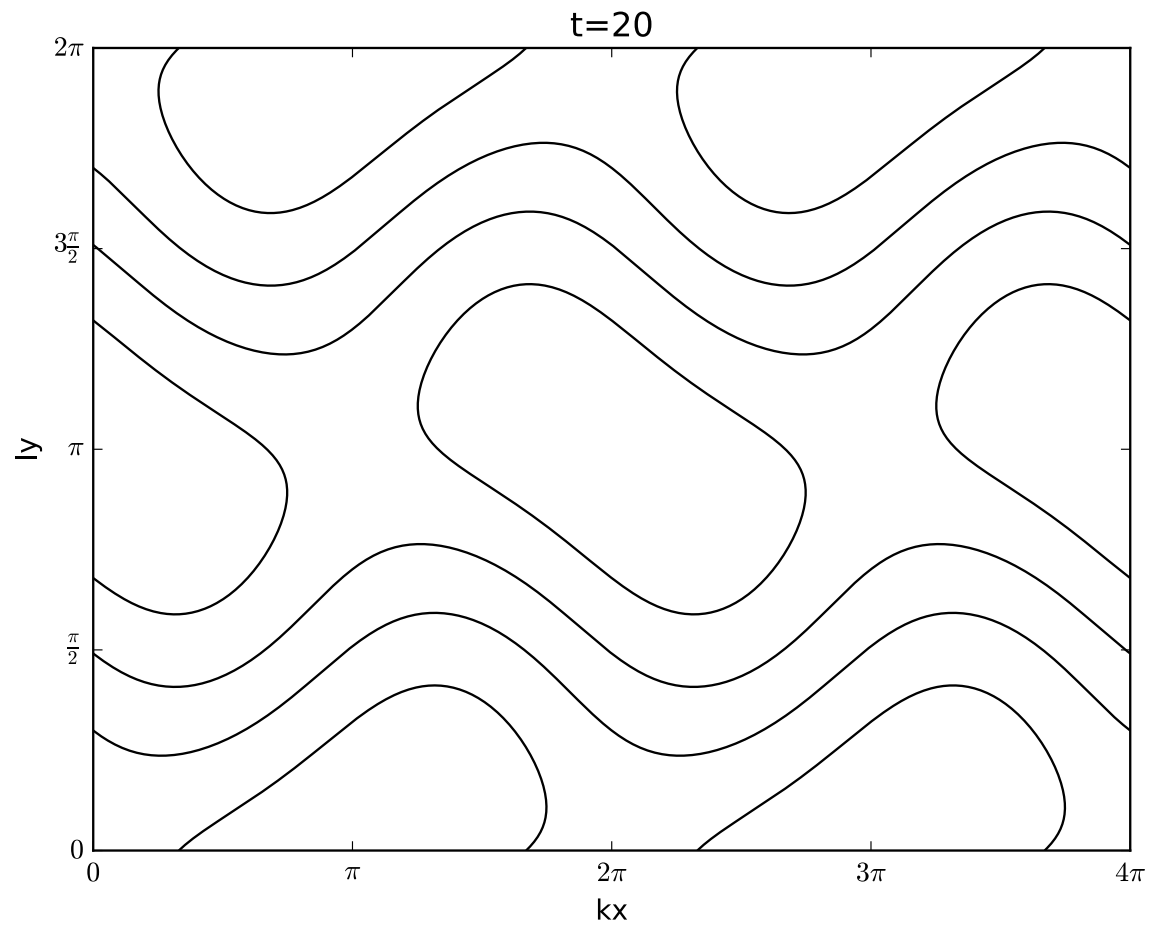
$$X_{20} = 0.24$$

$$X_{30} = 0.10$$

$$\alpha = 2$$

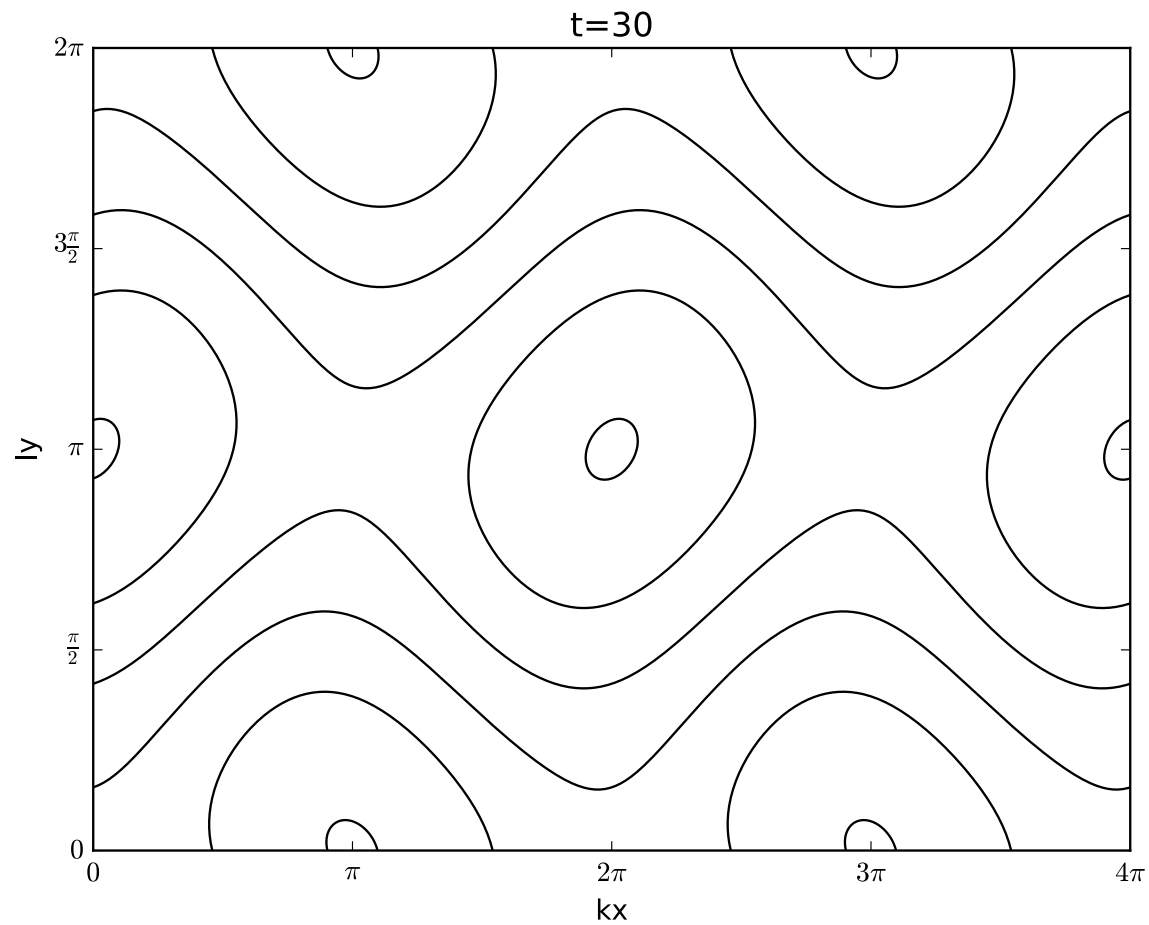
$$\beta = 0$$

# Analytical solution of the system



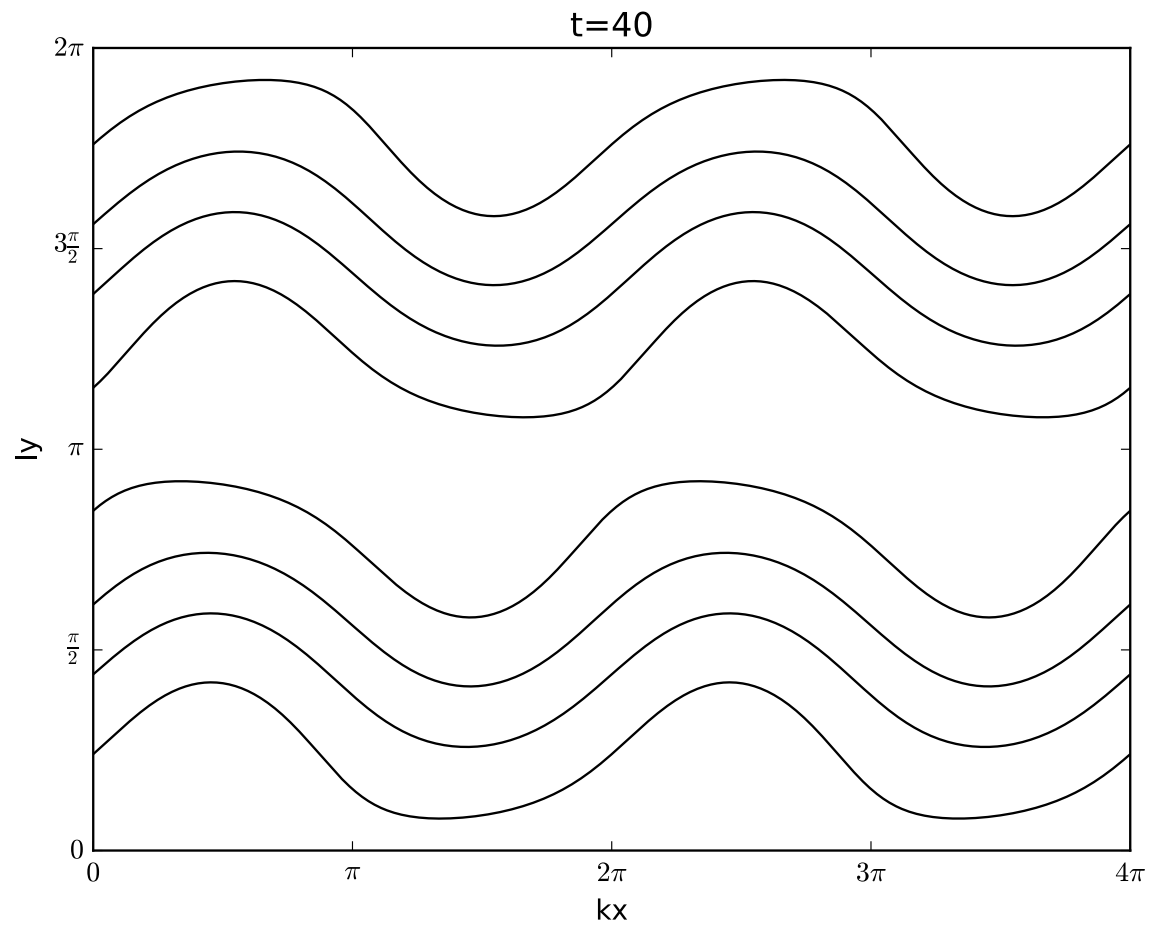
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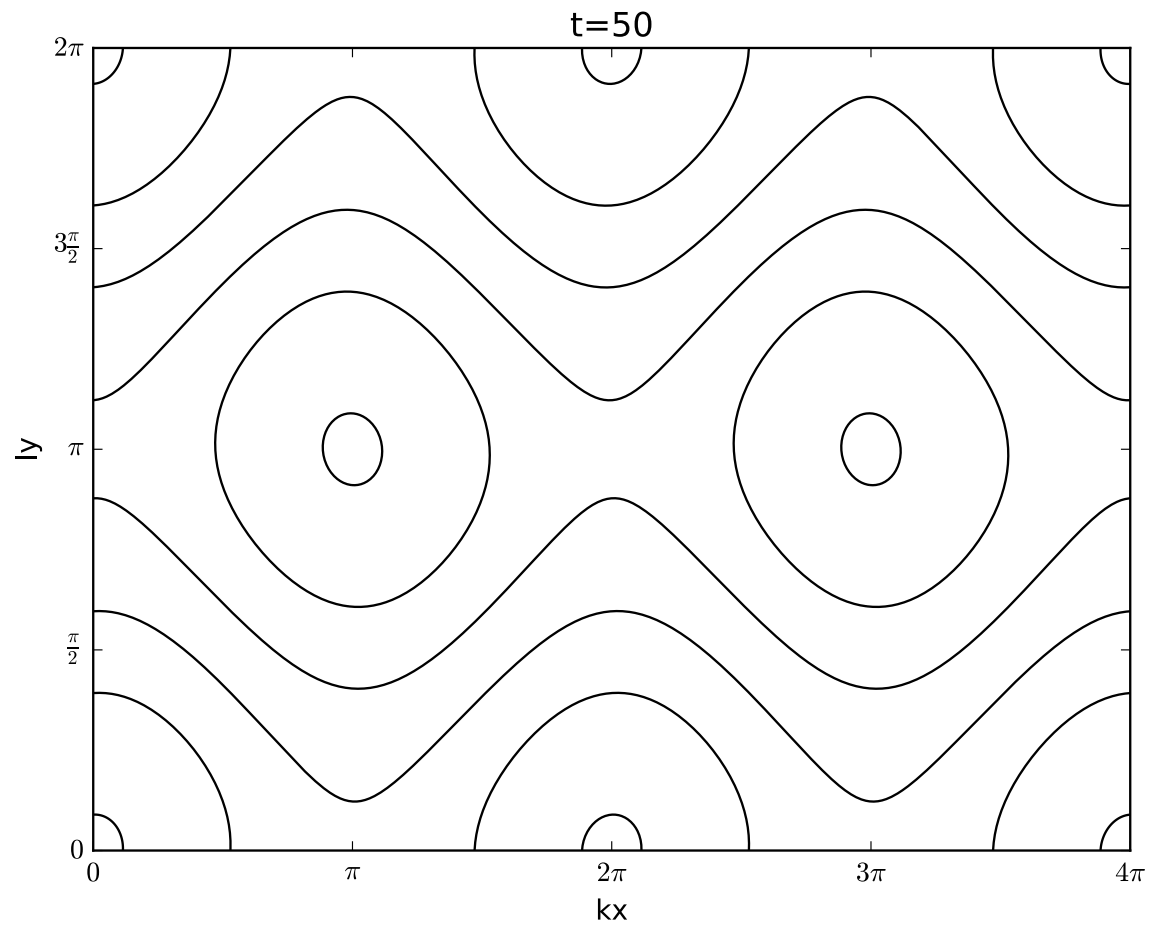
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# Analytical solution of the system



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$$\rho(\mathbf{X}, t) = \rho_0(\boldsymbol{\alpha}) \exp\left(-\int_0^t \chi(\mathbf{X}(\boldsymbol{\alpha}, t'), t') dt'\right)$$

## Solution of the Liouville equation

- Applying the general solution to the system obtained from the barotropic model yields:

$$\rho(X_1, X_2, X_3, t) = \rho_0(X_{10}(X_1, X_2, X_3, t), X_{20}(X_1, X_2, X_3, t), X_{30}(X_1, X_2, X_3, t))$$

## Solution of the Liouville equation

- The initial  $\rho$  is defined here as a three dimensional Gaussian distribution:

$$\rho_0(X_{10}, X_{20}, X_{30}) = k \exp \left[ - \left( \frac{(X_{10} - \mu_1)^2}{2\sigma^2} + \frac{(X_{20} - \mu_2)^2}{2\sigma^2} + \frac{(X_{30} - \mu_3)^2}{2\sigma^2} \right) \right]$$

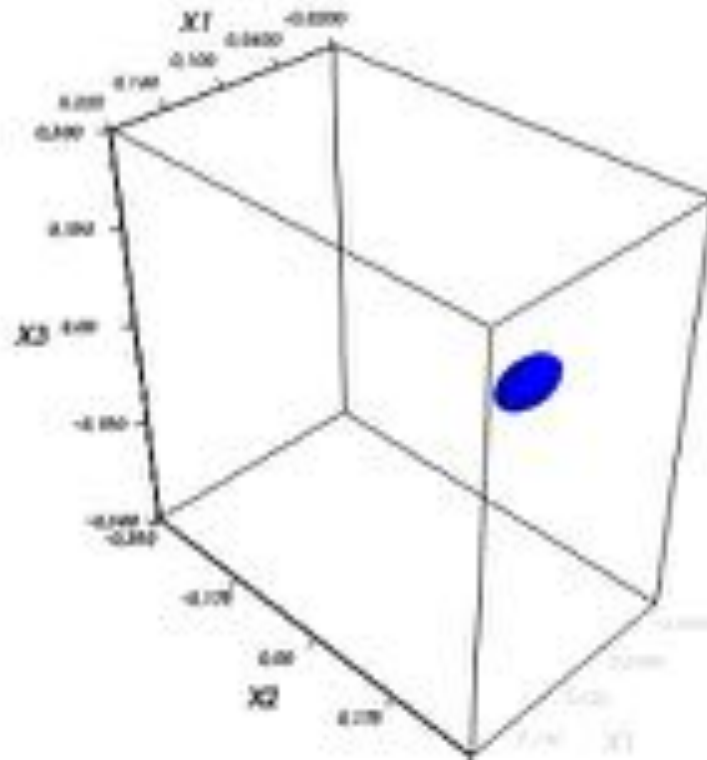
k: normalization constant

$\mu_k$ : means in each direction

$\sigma$ : standard deviation

- $\mu_1 = 0.12, \mu_2 = 0.24, \mu_3 = 0.1, \sigma = 0.01$

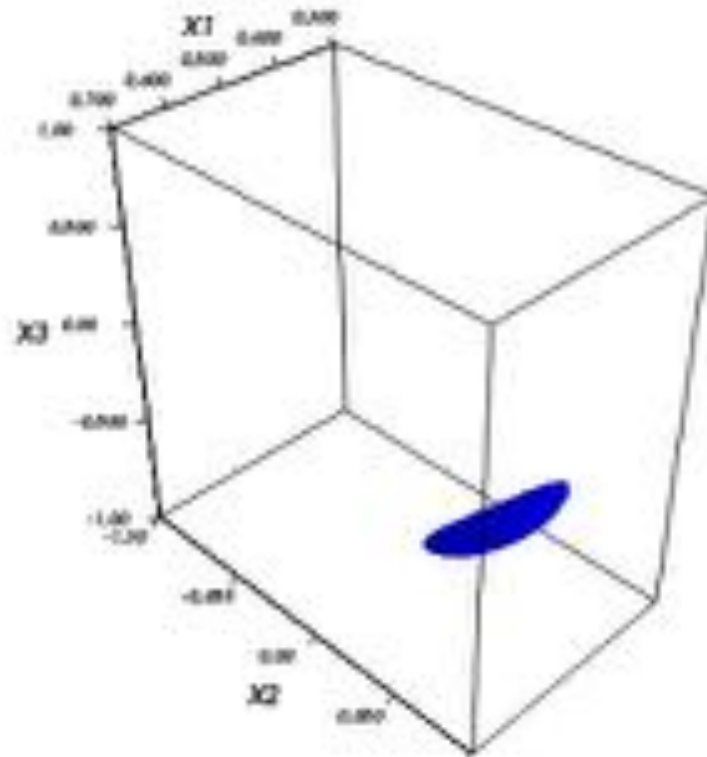
# Evolution of $\rho_1$ ( $\alpha = 2, \beta = 0$ )



$t = 0$

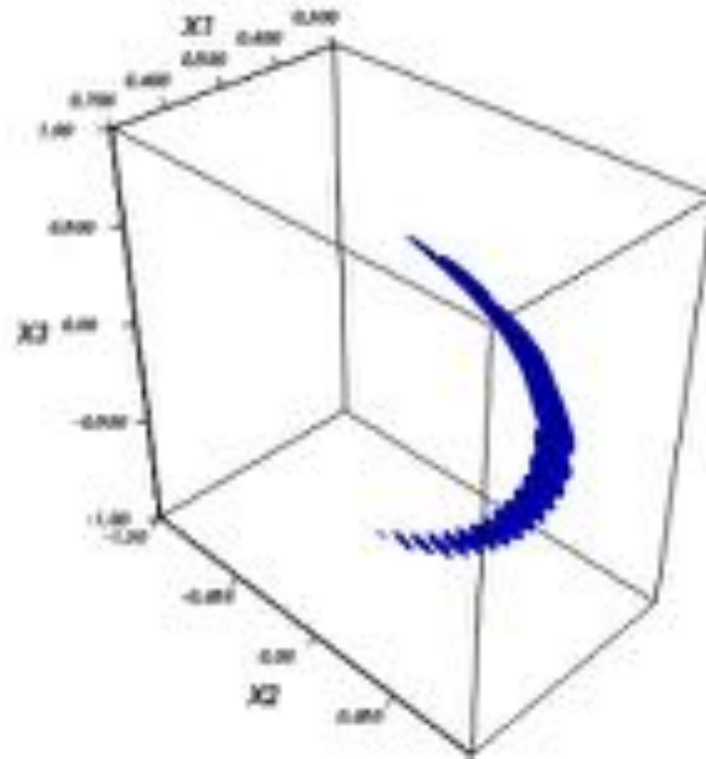
$$\alpha = \frac{k}{l}$$

Evolution of  $\rho_1$  ( $\alpha = 2, \beta = 0$ )



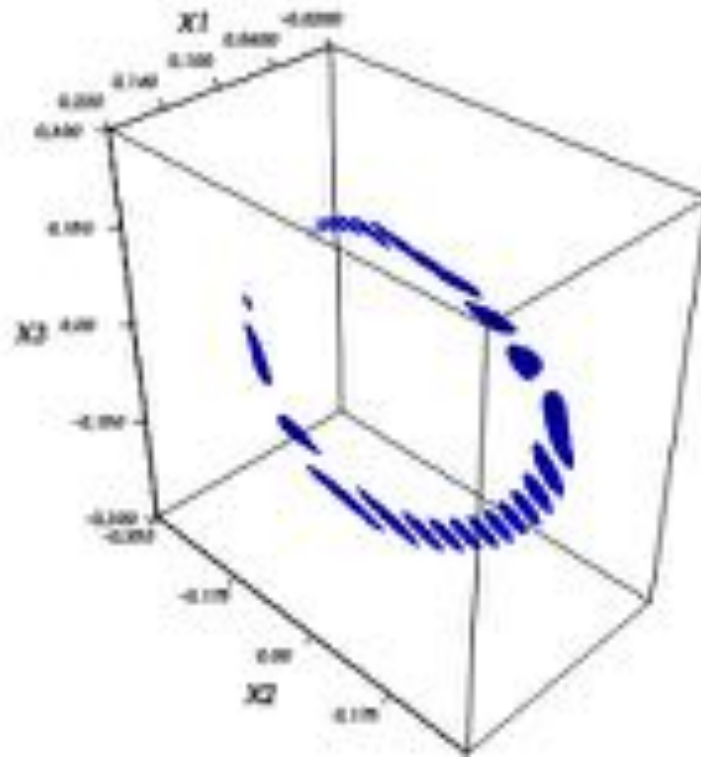
$t = 10$

Evolution of  $\rho_1$  ( $\alpha = 2, \beta = 0$ )



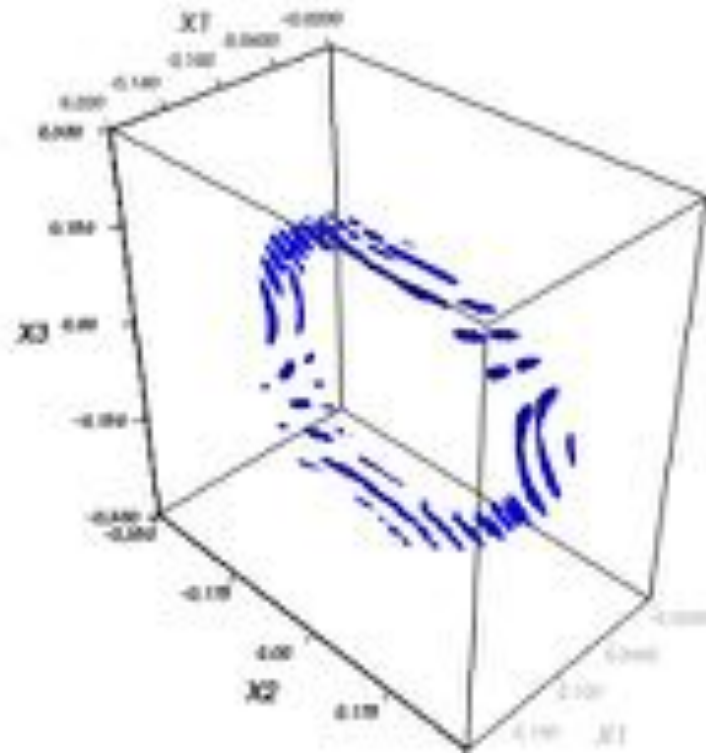
$t = 50$

Evolution of  $\rho_1$  ( $\alpha = 2, \beta = 0$ )



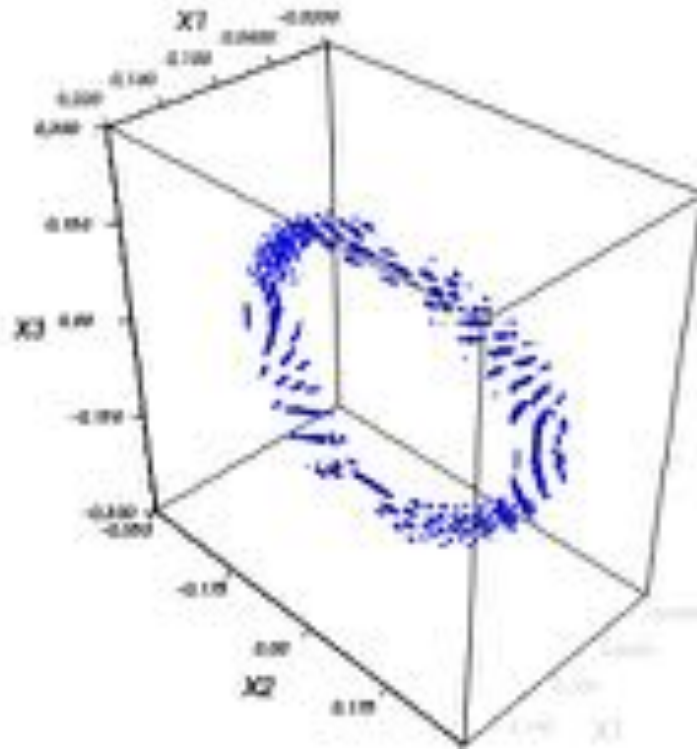
$t = 100$

Evolution of  $\rho_1$  ( $\alpha = 2, \beta = 0$ )



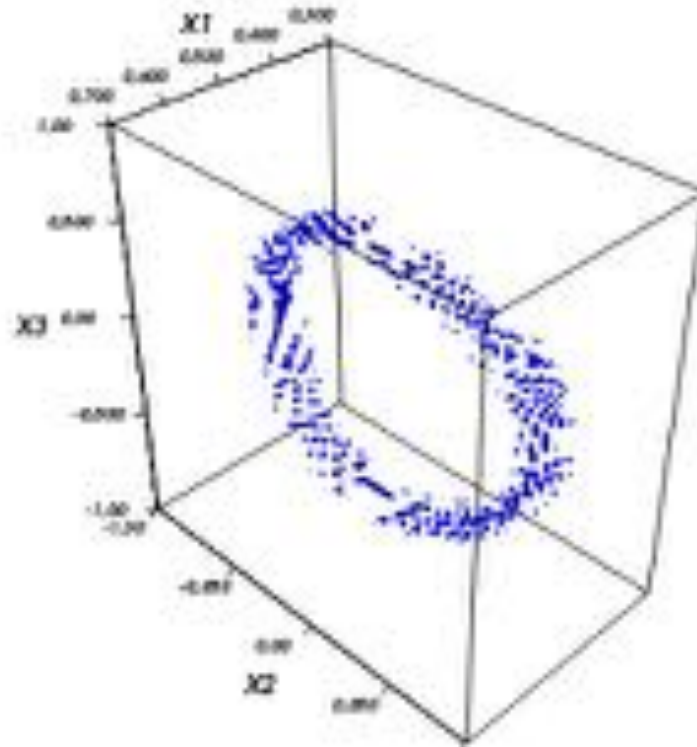
$t = 250$

Evolution of  $\rho_1$  ( $\alpha = 2, \beta = 0$ )



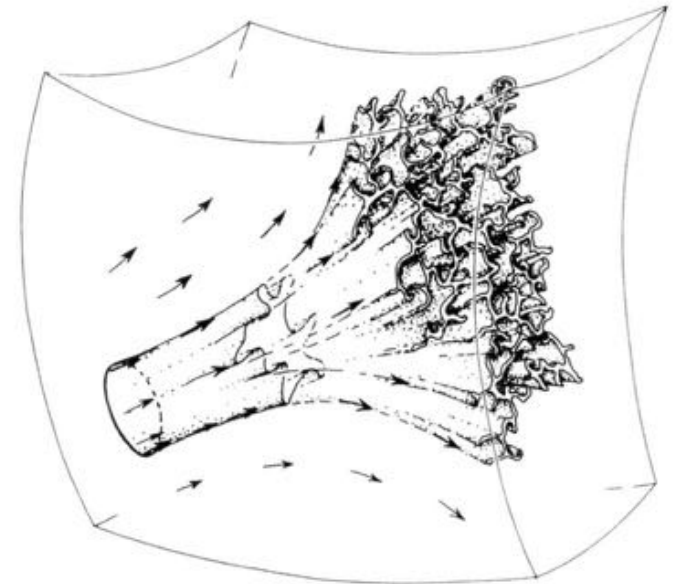
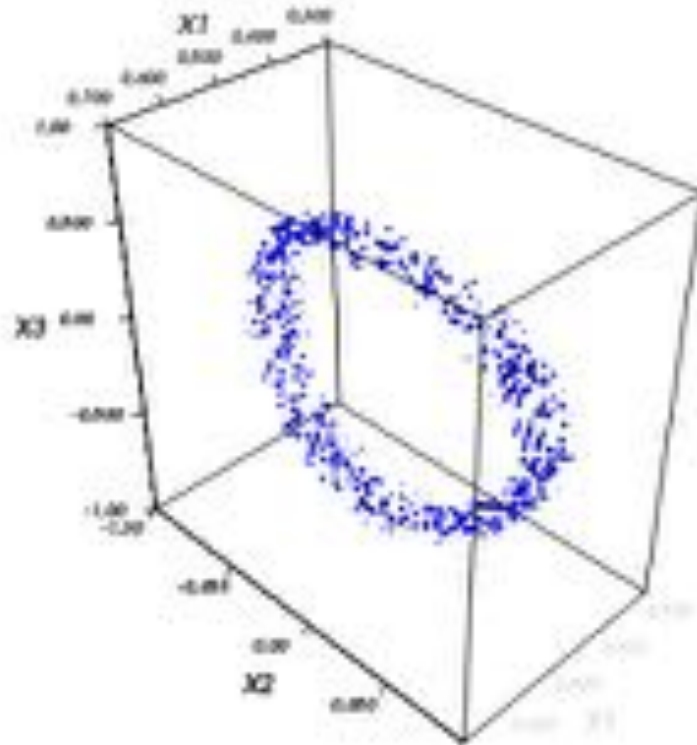
$t = 500$

Evolution of  $\rho_1$  ( $\alpha = 2, \beta = 0$ )



$t = 1000$

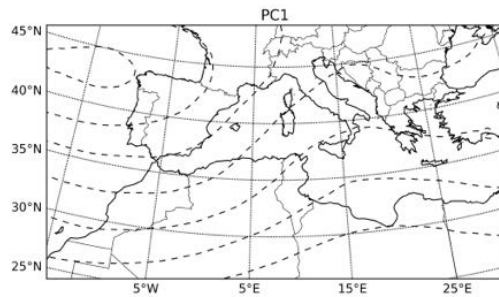
## Evolution of $\rho_1$ ( $\alpha = 2, \beta = 0$ )



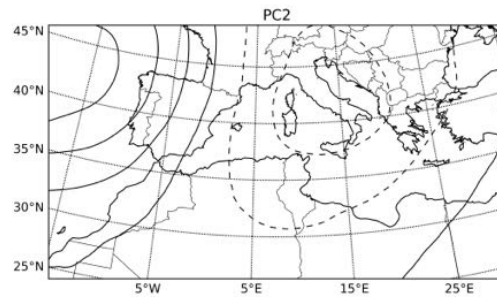
# Numerical solution of the Liouville equation for a barotropic model

- Solution of the Liouville equation for a full barotropic model in a 91x51 grid with a brute-force approach
- Total number of model integrations  $N_P^{N_D} O(10^{7000})$
- Principal Component Analysis is applied to reduce system dimensionality using ECMWF analysis data
- 5 PC are kept

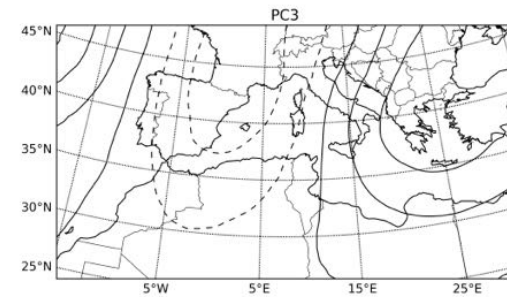
# Numerical solution of the Liouville equation for a barotropic model



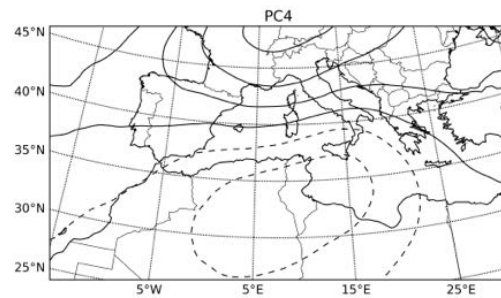
PC1



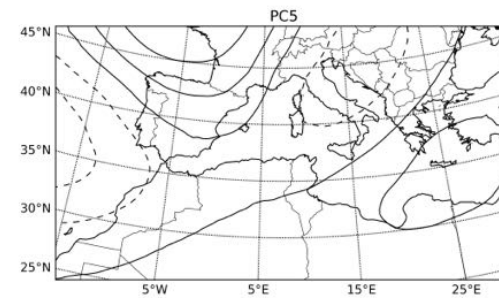
PC2



PC3



PC4



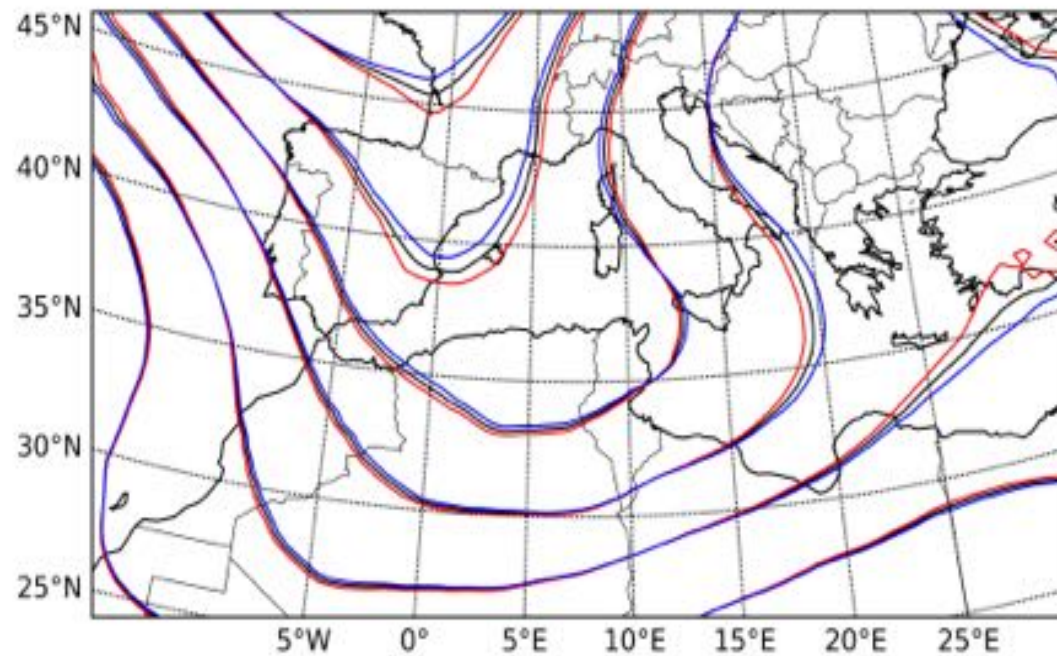
PC5

# Numerical solution of the Liouville equation for a barotropic model

- New phase space is 5-dimensional (5 PC retained)
- It is discretised in  $N_P = 75$  points in each dimension
- Initial PDF is a gaussian distribution
- Mean is obtained from the ERA5 500 hPa geopotential on 7<sup>th</sup> November 2014
- Liouville equation is solved with a lead time of 24h

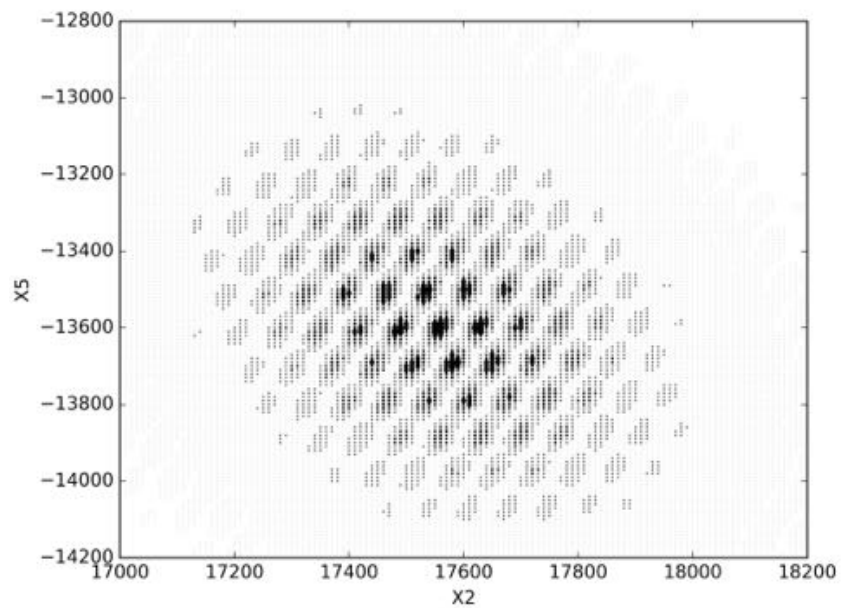
# Numerical solution of the Liouville equation for a barotropic model

- Initial conditions

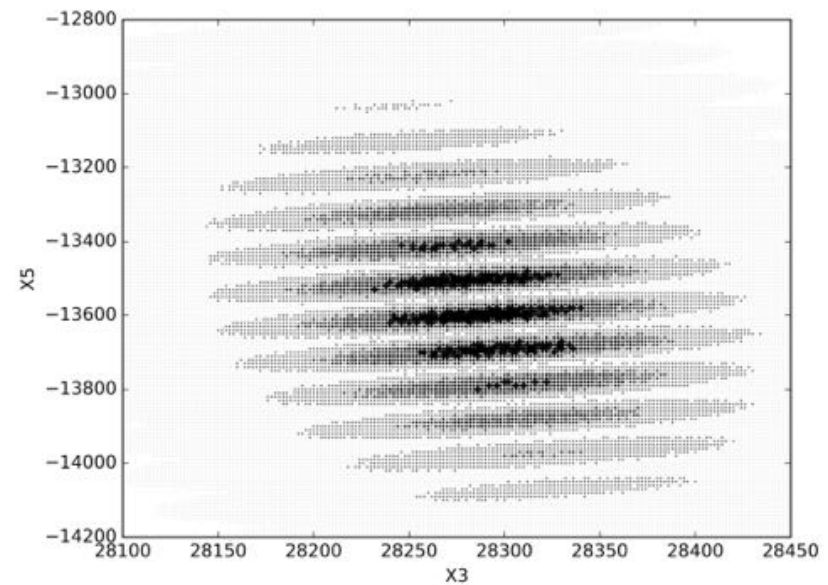


Mean + and -  $2\sigma$  in X3 for  $\rho_{bar1}$

# Results



X2-X5 plane



X3-X5 plane

## Part I conclusions

- The **Liouville equation** has been **solved analytically** for a low complexity system
- The **granularity** of the analytical solution for certain values of the parameters reveals a **serious predictability challenge**
- A **granular solution** is also identified in a more **realistic model**
- These results **challenge** most current **ensemble prediction** products that are based on compact PDF

## PART II: Tailored Bred Vectors

## Motivation and objectives

- One of the main problems of EPS is the underdispersion
- Extreme events may not be represented
- In order to improve the high resolution short-range forecast of extreme events, ensemble **spread** must be **controlled** (typically increased)

## New perturbation method: Tailored Bred vectors

- This method enables to **increase the ensemble size at no bred generation cost**
- Allows a **seamless scale representation** unlinking scales of forecast interest from bred generation strategy (and so recycling period)
- BVTEP are generated by combining different BV with different scales and amplitudes:

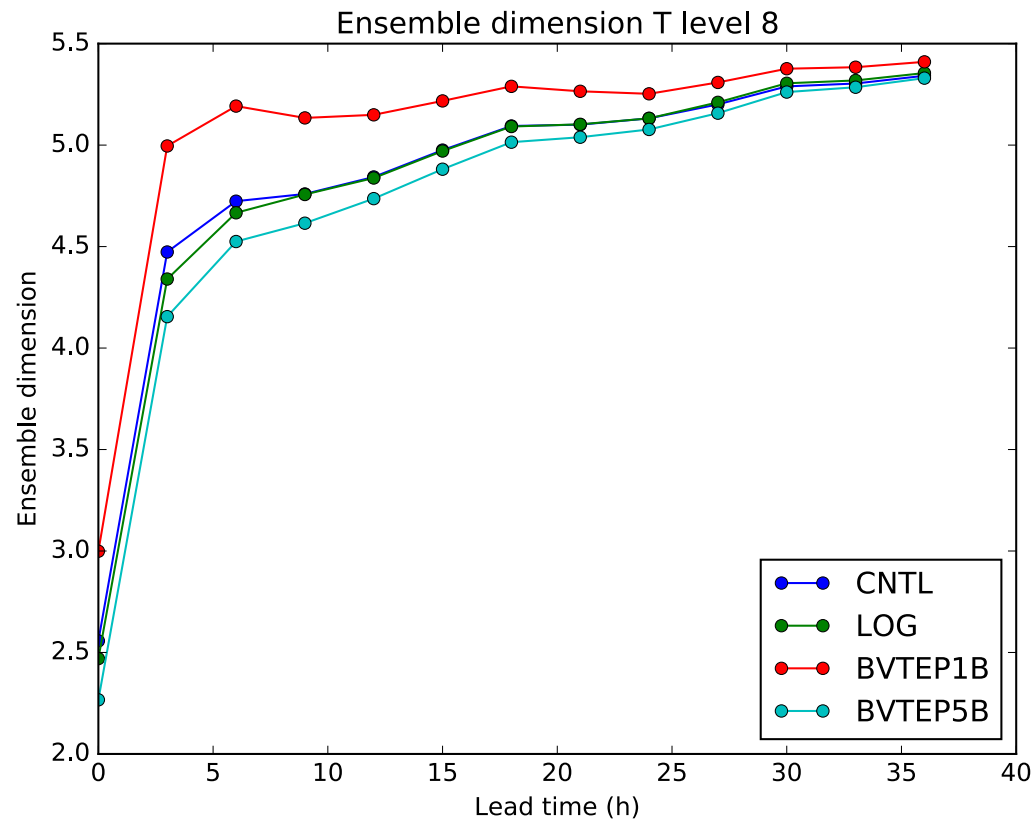
$$P = \sum_i \gamma_i \delta x_i^{1/\beta_i} \quad i: 1, \dots, n \text{ breds}$$

## Ensemble configurations

- **CNTL**: 5 ABV perturbations (1 bred per perturbation)
- **LOG**: 5 LBV perturbations (1 bred per perturbation)
- **BVTEP1B**: 5 ABV (1  $\omega$ -rescaled bred per perturbation)
- **BVTEP5B**: 5 ABV (Linear combination of up to 5  $\omega$ -rescaled bred per perturbation)

# Results

Variation of **ensemble dimension** with lead times for T at a model level close to 850 hPa



## Part II conclusions

- The **ensemble diversity** is **similar** for forecasts perturbed with **arithmetic** and **logarithmic** rescaled **bred vectors**
- Modifying the **scale** of the initial perturbations **increases ensemble diversity and skill**
- **Puzzling result:** better skill from particular case of general method...
- The methodology should be tested in a severe weather event

# Acknowledgments

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## ENSEMBLE DIMENSION

- **Ensemble dimension** is computed from the eigenvalues  $\mu_i$  of the covariance  $\mathbf{C}_{ij}$  matrix of a set of states:
- Perturbations covariance matrix:

$$\mathbf{C}_{ij} = \frac{\langle \mathbf{b}^{(i)}, \mathbf{b}^{(j)} \rangle}{L \|\mathbf{b}^{(i)}\|_2 \|\mathbf{b}^{(j)}\|_2} \langle \mathbf{b}^{(i)}, \mathbf{b}^{(j)} \rangle = \sum \mathbf{b}^{(i)}(\mathbf{x}) \mathbf{b}^{(j)}(\mathbf{x})$$

- Ensemble dimension

$$D = \frac{\left( \sum_{i=1}^k \sqrt{\mu_i} \right)^2}{\sum_{i=1}^k \mu_i}$$

D quantifies the number of independent vectors in the set (ensemble).